

Exploring the use of Unmanned Aerial Vehicles (UAVs) with the simplified “triangle” technique for Soil Water Content and Evaporative Fraction retrievals in a Mediterranean setting

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ABSTRACT

Information acquired from Unmanned Aerial Vehicles (UAVs) is frequently used nowadays in a variety of disciplines and research fields. The present study explores for the first time the combined use of UAVs with a newly proposed technique for estimating evaporative fraction (EF) and surface soil moisture (SSM). The investigation is performed in a typical Mediterranean setting, a citrus field with flat topography divided in two plots with different irrigation schemes, in Sicily, Italy, at which ground data acquired during an extensive field campaign in July 2019. Reasonable estimates of both EF and surface wetness were produced, with patterns in agreement to vegetation cover fragmentation, topography, and other site-specific characteristics. Validation shows average error of 0.053 for EF and of 0.040 cm³ cm⁻³ for SSM. The results are comparable or better to those reported in analogous studies performed in similar areas. This implies that the investigated approach performs well under the semi-arid conditions characterising the experimental set up. To our knowledge, this study represents the first evaluation of the combined use of the “simplified triangle” with very high-resolution UAV imagery. As such, the findings are of significance regarding the potential future use of the “simplified triangle” approach particularly with very fine resolution imagery such as that provided by UAV for mapping and monitoring EF and SSM in agricultural and natural ecosystems.

KEYWORDS: earth observation, unmanned aerial vehicles, surface soil moisture, evaporative fraction, simplified triangle, surface temperature/vegetation index

1. Introduction

The natural processes taking place on the Earth’s surface control the energy and mass exchanges between land and atmosphere and are key drivers of the Earth’s system (North et

47 al., 2015; Gerken et al., 2019). Today, particularly so in light of climate change and concerns
48 related global food and water security, an improved understanding of land-atmosphere
49 interactions is a topic of urgent importance (Ireland et al., 2015; Deng et al., 2019). In this
50 context, obtaining accurate information on the spatial and temporal variability of land surface
51 parameters such as evaporative fraction, EF (defined as the ratio of instantaneous latent heat
52 flux (LE) to net radiation (R_n) and surface soil moisture (SSM) is of primary interest for several
53 environmental applications and research investigations (Jung et al., 2011; Srivastava et al.,
54 2019). This is due to the influence of these parameters on key physical processes and
55 feedback loops of the Earth system (Nutini et al., 2104; Srivastava et al., 2015; Amani et al.,
56 2016). Accurate information on their spatiotemporal variability, particularly at fine spatial and
57 temporal resolution, can provide valuable information in research studies and practical
58 applications linked to ecosystem processes, plant water requirements and water resources
59 management (Shi et al., 2014; Minacapilli et al., 2015; Deng et al., 2019; Yang et al., 2020).

60 Despite their significance, it is quite difficult to quantify EF and SSM on a routine basis over
61 large geographical regions using ground instrumentation. The main reasons include the large
62 spatiotemporal variability of these parameters (Bao et al., 2018). Earth Observation (EO)
63 presents a suitable alternative to ground observations for deriving SSM and/or EF over large
64 regions and diverse geographical scales (Tian et al., 2014). A variety of approaches have
65 been proposed for this purpose, ranging from semi-empirical to physically-based ones (see
66 Petropoulos et al., 2015; 2018). Those approaches are characterised by different degree of
67 complexity, input parameters requirements and retrieval accuracy.

68 A specific group of EO-based techniques commonly termed as surface temperature (T_s) and
69 vegetation index (VI) methods (T_s /VI), has shown an excellent promise at deriving spatially
70 explicit maps of sensible and latent heat fluxes (H, LE) and/or SSM. These methods utilise
71 optical (visible and infrared - VNIR) and thermal infrared (TIR) EO data and are based on
72 physical relationships between the satellite-derived T_s and a VI, the latter being associated to
73 the existent degree of vegetation (Zhang et al., 2014; Capodici et al., 2020). If these
74 parameters are in a scatter plot, provided that there is a full variability in VI, a
75 triangular/trapezoidal shape similar to that shown in **Figure 1** emerges. This shape,
76 characterised by the physical boundaries also shown in **Figure 1**, results from the T_s
77 sensitivity to water content, which increases as a function of the proportion of bare soil
78 exposed. The biophysical properties included in this T_s /VI domain are well-documented
79 (Gillies et al., 1997; Chauhan et al., 2003; Maltese et al., 2015; Wang et al., 2018; Cui et al.,
80 2020). Detailed descriptions of these properties, including the key parameters affecting the
81 T_s /VI scatterplot shape, are summarised in Petropoulos et al. (2009) and Petropoulos et al.
82 (2018). Tang et al. (2017) introduced the End-member-based Soil and Vegetation Energy
83 Partitioning model (ESVEP), a two-source approach for estimating land surface
84 evapotranspiration (ET)) for which two dry edges could be considered in the case of a root
85 zone water stress occurs. It is based on the consideration that soil evaporation primarily draws
86 water from the upper soil layer, whereas, transpiration exploits water from the root zone. The
87 temporal response of soil water content of the upper soil and root zone in the framework of the
88 ET process is therefore different: the dynamic of the soil water content is more rapid in the
89 upper layer; it is slower in the root zone.

90 Recently, Carlson & Petropoulos (2019) proposed a T_s /VI technique for estimating both EF
91 and SSM, which they named “simplified triangle”. This approach is essentially a variant of the
92 so-called “triangle” technique (Carlson, 2007) and does not require for its implementation a

93 land biosphere model or any other ancillary data. These characteristics make this approach
94 easy to apply in comparison to other T_s/VI methods. Furthermore, being dependent on a small
95 number of easily computed EO-based parameters, it becomes a very attractive choice for
96 potential operational use. Fuzzo et al. (2019) demonstrated how this newly introduced
97 approach can be coupled with a crop prediction and a climatological water balance model in
98 soybean yield prediction using MODIS data. However, as the technique is recent, studies
99 validating its performance in different environments and with a range of EO instruments are
100 scarce.

101

102 [Please put Figure 1 around here]

103

104 To our knowledge, this newly proposed technique has not been implemented on and verified
105 for unmanned aerial vehicles (UAVs) data yet. UAV platforms with on-board visible/near
106 infrared and thermal sensors have very important advantages over satellite EO platforms,
107 such as user flexibility to select the target area and the frequency of data acquisition (Dawson
108 et al., 2019; Liu et al., 2020). Therefore, this technique implementation with UAVs would be
109 indisputably of key importance, as it would inform on its potential usefulness in a broad
110 spectrum of practical applications and research purposes alike.

111 In this context, this study aims at exploring for the first time the combined use of the “simplified
112 triangle” with very high spatial resolution UAV data, to predict the spatio-temporal variability of
113 both EF and SSM. For this purpose UAV, ground truthing and ancillary data acquired during a
114 field campaign that took place in July 2019 at one experimental site in Sicily, Italy, are
115 employed. The experimental set up description is provided in Section 2, whereas the
116 “simplified” technique implementation with the UAV data is made available in Section 3,
117 followed by the results and the related discussion which are described in Sections 4, and 5,
118 respectively.

119

120 **2 Materials**

121 **2.1 Study site**

122 The study site is a citrus orchard field (*C. reticulata* Blanco, cv. Tardivo di Ciaculli) located in
123 the neighbourhood of Palermo, Italy (38° 4'53.4"N, 13° 25' 8.2"E). The site contains 30 year
124 old tangerine trees planted at a regular spacing of 5.0 m × 5.0 m (plant density of 400 plants
125 per ha) and irrigated with a subsurface drip system. The area is in a typical eastern
126 Mediterranean semi-arid environment. The study area has flat topography with elevation
127 between 30 and 35 m above sea level, and slopes ranging from 1% to 4%.

128 To differentiate irrigation management, the field has been divided into two plots of about 4,000
129 m² each, as shown in **Figure 2**. The first plot was maintained under full irrigation (FI), whereas
130 the second under deficit irrigation (DI) applied throughout phase II of fruit growth (from 1 July
131 2019 to 20 August 2019). Each plot was, in turn, divided into four sub-plots differentiated for
132 the anti-root agents introduced into the emitters during the manufacturing process, but not for
133 the irrigation management, nor for the emitters' hydraulic performance. The subsurface drip
134 system is characterized by two lateral pipes per plant row, installed at a distance of 1.1 m from
135 the trees and buried at a depth of 0.30 m. In each lateral pipe, self-compensating emitters

136 were installed with half-meter spacing between them, nominal flow rate of 2.3 h⁻¹ and
137 operating pressure of 150 kPa. A disc filter, an electric control valve, a relief valve, a pressure
138 gauge, and a flow meter completed each sub-plot irrigation unit.

139 The experimental setup is equipped with a WatchDog 2000 weather station (Spectrum
140 Technologies, Inc.), including sensors for relative air humidity, wind speed and direction, air
141 temperature, solar radiation, and rainfall, as well as eight "drill & drop" frequency domain
142 reflectometry sensors (Sentek Pty Ltd, Stepney, Australia) to monitor soil water content,
143 installed on a central tree of each sub-plot, 0.30 m away from the closest emitter. All the
144 sensors were interfaced with a communications board that uses the cellular 3G data network
145 for internet connection using the MODBUS RTU protocol to transfer and save the data into a
146 MySQL database operated by AgriNET/Tuctronics which is accessible from the web. The
147 system allows the download of weather variables, soil water content (SWC) and temperatures
148 (T) in the root zone, at 10 cm intervals from the first 5 cm of the soil layer down to a depth of
149 0.6 or 1.2 m. The Scholander chamber (Scholander et al., 1965) was used to follow the
150 temporal dynamic of predawn and midday stem water potential, whereas a couple of Granier
151 thermal dissipation probes (Granier, 1985) was installed in four trees to monitor sap flow
152 during the irrigation season.

153 In addition, an eddy covariance flux tower was set up in the orchard in February 2019 to
154 measure the turbulent fluxes (sensible, H, and latent, LE, heat fluxes) and a four-component
155 net radiometer was used to measure net radiation (R_n) individual components. A CNR1 four
156 component Net Radiometer was installed at 3.1 m a.g.l, while an InfraRed Gas Analyzer IRGA
157 LI7500 (manufactured by LI-COR, Inc.) and a CSAT3 Three Dimensional Sonic Anemometer
158 anemometer (manufactured by Kipp & Zonen B.V.) were installed slightly above, at 3.5 m
159 above ground level (a.g.l.), , i.e., approximately 55 and 95 cm above the vegetation canopy.
160 All the data were processed at 30 minutes interval. The footprint flux tower was calculated
161 according to Schuepp et al. (1990) at 70% of the fluxes.

162 [Please put Figure 2 around here]

163

164 2.2 Data Acquisition & Pre-processing

165 2.2.1 Data Acquisition

166 The fieldwork for this study was carried out on July 2019. A series of spatial and ancillary data
167 was acquired on 30 July 2019 as part of the field campaign that was conducted in order to
168 support the study implementation. In particular:

169 ➤ **Global Navigation Satellite System (GNSS) Survey.** Nine black and white control
170 targets, and the same number of aluminium targets were distributed on a regular grid
171 to cover the whole study area.

172 The coordinates of the targets were measured by a NRTK survey using a Topcon
173 Hiper V receiver (both Global Positioning System (GPS) and Glonass constellations).
174 A UNIPA (University of Palermo) GNSS Cross-origin resource sharing (CORS)
175 network encompassing 8 permanent stations, 2 of them installed on two University
176 buildings in Palermo and Agrigento and 6 at other public institutions of the Sicilian
177 territory was employed for Network real-time kinematic (NRTK) positioning. The
178 network covers about 7400 km² western Sicily. The GNSS CORS Network project was
179 carried out with the technical collaboration of Topcon Italy (that supported the scientific

180 research with GNSS receivers and antennae), in the framework of developing a
181 network for technical (real-time) and scientific (post-processing) use. The CORSs
182 included in the Topcon Netgeo GNSS network. Since 2013 the data retrieved from
183 UNIPA GNSS CORS network have been used for the computation of the RDN2 (*Rete*
184 *Dinamica Nazionale 2*) which provides the WGS84 datum for Italy in the European
185 Permanent Network (EPN subnetwork). UNIPA GNSS CORS network has received
186 the scientific acknowledgment through many experiments in various application fields
187 (Catania et al. 2020; Angrisano et al. 2020, Kenyeres *et al* 2019, Pipitone et al. 2018,
188 Dardanelli et al. 2015, Dardanelli et al. 2014, and Dardanelli and Carella, 2013). Since
189 2013 the postprocessing RINEX (Receiver INdependent EXchange) data have been
190 made available for the evaluation of the national reference framework by the IGMI
191 (Italian cartographic military institute) and for technical researches able to investigate
192 the horizontal and vertical velocity map in Italy (Maseroli, 2015). NRTK positioning was
193 carried out using the hardware and software infrastructure of the permanent Netgeo-
194 Topcon Italy network framed in the reference system ETRF2000 (powered by UNIPA
195 GNSS CORS) and in particular via the VRS (Virtual Reference Station) stream. Data
196 availability and geodetic framework are described in Dardanelli *et al.* (2020). The
197 processing of GNSS data acquired to allow an accurate orthorectification of
198 multispectral and thermal images was carried out by Meridiana software ver. 2020.

- 199 ➤ **Proximity sensing images.** Multispectral images were acquired using a NT8 contras
200 octocopter carrying a RikolaDT-17 Fabry-Pérot camera (manufactured by Rikola Ltd).
201 The multispectral camera has a 36.5° Field of View. It was set-up to acquire images in
202 9 spectral bands with a 10 nm bandwidth. Central wavelengths were 460.43, 480.01,
203 545.28, 640.45, 660.21, 700.40, 725.09, 749.51 and 795.53 nm. At a flight altitude of
204 50m above ground (a.g.l.), the average Ground Sampling Distance (GSD) was 3 cm.
205 Thermal images were acquired almost simultaneously to the multispectral images,
206 using a DJI Mavic 2 Enterprise Dual quadcopter carrying on-board a FLIR Lepton®
207 (manufactured by FLIR® Systems, Inc) acquiring in the longwave infrared spectral
208 range (from 8 to 14 μm), with a thermal sensitivity lower than 50 mK (0.050 °C). The
209 average GSD was 3.46 cm. All the images were resampled at 4 cm spatial resolution
210 using a pixel aggregate resampling method.
- 211 ➤ **Spectroradiometric measurements.** Four reference targets, ranging in a greyscale
212 from black to white were also positioned to allow the spectral reflectance calibration by
213 means of a field spectroradiometer. The employed ASD FieldSpec®FR
214 spectroradiometer (Analytical Spectral Device, ASD, Inc.) measured the full solar
215 spectrum (between 350 and 2500 nm) with no fore optic attached.
- 216 ➤ **Thermographs.** Ground measurements of surface temperature (T_s) were carried out
217 at noon using a handheld FLIR SC660 (FLIR® Systems, Inc.) characterized by a
218 sensitivity lower than 30 mK.

220 2.2.1 Pre-processing

221 Following the data acquisition, standard pre-processing steps were applied. To orthorectify the
222 multispectral and thermal images, a standard photogrammetric/SfM approach (e.g., Harwin
223 and Lucieer, 2012) was applied via Pix4D mapper (by Pix4D Inc.). A Topcon Hiper V receiver
224 (both GPS and GNSS Connectivity) was employed to acquire ground control points for the
225 orthorectification. The average position dilution of precision (PDOP) and the geometric dilution
226 of precision (GDOP) were 1.8 and 2.0, respectively. The control targets were positioned with

227 average planimetric and altimetric accuracy of ± 2 cm that can be considered within acceptable
228 geometrical configuration limits to orthorectify the UAV images, considering that these latter
229 are characterized by a spatial resolution of 4 cm once orthorectified. Images acquired in the
230 visible and near infrared were calibrated to ground reflectance implementing the empirical line
231 technique (Karpouzli and Malthus, 2003), which allows the simultaneous correction of the
232 atmospheric influence. Similarly, TIR images were calibrated into surface radiometric
233 temperatures by means of a linear regression with at ground thermographs and an emissivity
234 map of the soil vegetation system (Negm et al., 2017). The spatial distribution of emissivity
235 was calculated according to Valor and Caselles (1996). Given the spatial resolution of the
236 images (about 10^{-2} m) compared to the spacing of the trees (about 5 m) we did not consider
237 the cavity effect. We assume the emissivity values for bare soil and densely vegetated ground
238 to be equal to 0.97 and 0.99, respectively, as reported in Sobrino et al. (2004). **Figure 3**
239 illustrates the Normalized Difference Vegetation Index (NDVI) and of Surface Temperature
240 (T_s) final products upon completion of all pre-processing steps.

241 [Please put figure 3 around here]

242

243 3 Methods

244 3.1 Simplified Triangle Method

245 A comprehensive account of the “simplified “triangle technique implementation is available in
246 Carlson and Petropoulos (2019). Briefly, the method allows the retrievals of two parameters,
247 the soil water availability (M_o) and EF. M_o represents surface wetness in the bare soil surface
248 (top few millimetres of it) and it is computed from the ratio between the actual soil/vegetation
249 system evapotranspiration ET and potential evapotranspiration (ET/ET_p). M_o is also equated
250 to SSM by multiplying M_o with the soil’s field capacity. On the other hand, EF is defined as the
251 ratio between latent heat flux (LE) and net radiation (R_n).

252 EF and M_o are obtained from the T_s/VI feature space. The scatterplot is constructed by plotting
253 the T_s versus fractional vegetation (F_r), where the latter is computed from the NDVI (see
254 Equation (1) below) and its corresponding range of variability, as proposed by Carlson (2007).
255 Upon completion of this step, a number of parameters need to be determined, namely: (a) the
256 NDVI values for bare soil and dense vegetation (respectively, $NDVI_o$ and $NDVI_s$), and (b) the
257 highest value of T_s (T_s [max]) which is characteristic of dry/bare soil pixels, as well as the
258 minimum value of T_s (T_s [min]).

259 $NDVI_o$, $NDVI_s$, T_{max} and T_{min} are used to specify the T_s/VI feature space boundaries and to
260 constrain the solution for EF and M_o . $NDVI_s$ and T_{min} , represent dense vegetation and define
261 the lower left (wet) vertex of the triangle, i.e. the so-called ‘wet edge’ or ‘cold edge’ (see
262 **Figure 4**). The wet edge corresponds to M_o and EF values equal to 1.0. Similarly, $NDVI_o$ and
263 T_{max} define the lower right vertex of the triangle, the so-called ‘dry edge’ or ‘warm edge’ (also
264 shown also in **Figure 4**). These points characterize the soil dryness boundary with $M_o = 0$ and
265 covers the area from T_{max} and $NDVI_o$ to $NDVI_s$, which, for a triangle with a distinct upper
266 vertex, occurs at T_{min} . Even though $M_o = 0$ along the “dry edge”, along the dry edge EF itself is
267 non-zero apart from the triangle’s lower right vertex. The next step in the technique
268 implementation includes the scaling of T_s to T^* (by applying Equation (2) below), which ranges
269 between zero to one.

270 At this stage two central hypotheses are made. The first is that when vegetation is at wilting
 271 point transpiration is always equal to the potential transpiration, as generally assumed in
 272 nearly all T_s/VI approaches (e.g., Jiang and Islam 2003). The second hypothesis is related to
 273 the relationship between EF and M_o within the T_s/VI domain, which is assumed to be linear.

274 [Please put Figure 4 around here]

275

276 Thus, on the basis of the assumptions above, M_o is defined as the ratio between the lengths
 277 “a” and “d”. Both these lengths depend on T^* and F_r . For conditions where a pixel comprises
 278 of both areas of vegetation and bare soil, the canopy EF is taken as the weighted value of EF
 279 for the vegetation fraction of the pixel ($EF_{veg} = 1$, by definition). As such, both M_o and EF are
 280 computed for all pixels contained in the T^*/F_r domain from the implementation of Equations (3)
 281 and (4) shown below.

282

$$F_r = \left\{ \frac{(NDVI) - (NDVI_0)}{(NDVI_s) - (NDVI_0)} \right\}^2 \quad (1)$$

$$T^* = \{T - T_{min}\} / \{(T_{max} - T_{min})\} \quad (2)$$

$$M_o = 1 - T^*(pixel) / T^*(dry\ edge) \quad (3)$$

$$EF = (EF_{soil})(1 - F_r) + F_r (EF_{veg}) = M_o(1 - F_r) + F_r \quad (4)$$

283 In the above, EF_{soil} refers to the ratio between soil evaporation and net radiation. T (pixel) is
 284 the scaled surface temperature T^* for a given pixel within the scatterplot and T (dry edge) is
 285 the value of T^* at the dry edge of the triangle. In this study, the values for the temperature limits
 286 were $T_{min} = 19.40$ °C and $T_{max} = 73.27$ °C, whereas for NDVI were $NDVI_0 = 0$ and $NDVI_s = 1$.
 287 Noticeably that fully vegetated pixels exhibit a variability in T^* of 0.25 conferring to the $T^* - F_r$
 288 scatterplot a trapezoidal shape. The variability in T^* could be attributed to the very high spatial
 289 resolution achieved by UAV which allows to record the surface temperatures of the single
 290 leaves of the same canopy. In particular, the variability in T^* is attributed to the different
 291 exposure to the direct solar radiation of the single leaves which controls i) directly, the
 292 individual leaf warming up; ii) indirectly, the leaf transpiration.

293 The implementation of the steps summarized above to the pre-processed UAV data resulted
 294 in the scatterplots of NDVI vs T_s and of computed F_r vs T^* shown in **Figure 5**. The spatial
 295 maps of F_r and T^* are also shown in this figure.

296 [Please put Figure 5 around here]

297

298 3.2 Statistical Analysis

299 Evaluation of the predicted SSM and EF included at first a visual inspection of the
 300 spatiotemporal variability of the derived maps. Next, the main validation approach involved
 301 comparisons at pixel level between the predicted and measured parameters. The statistical
 302 scores computed that quantify the agreement between predictions and observations are
 303 summarised in Table 1. These statistical measures have already been used in similar past
 304 verification exercises (e.g. Nutini et al., 2014; Piles et al., 2016; Amani et al., 2016; Xu et al.,
 305 2018, Wang et al., 2018).

306 [Please put Table 1 around here]

307

308 **4 Results**

309 4.1 Visual Comparisons

310 The EF and SSM maps and their corresponding histograms obtained from the UAV data and
311 the “simplified triangle” technique are illustrated in **Figure 6**. The first step of the analysis
312 included a visual inspection of the spatial variability of the derived parameters. As can be
313 observed, both EF and SSM maps exhibited a sensible range of values as well as reasonable
314 spatial variability. Clearly, the spatial variability is in agreement with the changes in land
315 use/cover, as well as with the derived F_r and T_s maps based on the UAV data that were
316 presented in **Figure 5**. Both EF and SSM predicted by the “simplified” triangle are spatially
317 consistent with the soil/vegetation cover patterns and variability: in particular, high values of
318 both variables correspond to the vegetated areas of the image, whereas low values appear in
319 areas of bare soil.

320 To further illustrate the above observation, it was further investigated the variability of the
321 derived parameters separately for the bare soil and the partially or fully vegetated components
322 (see **Figure 7**). As evidenced in the maps shown in **Figure 7** (and their associated
323 histograms), the variability of the examined parameters is largely explained by the spatial
324 variability in the land surface fragmentation. It is evident from the visual comparisons of bare
325 soil and vegetation maps and histograms, that the variability of the vegetation for both EF and
326 SSM is significantly higher in bare soil. From these figures it is shown that the EF and SSM for
327 vegetation are predominantly above 0.9 EF and 0.2 SSM. Bare soil presents higher variability,
328 but the highest frequencies (especially for SSM) are close to 0.26.

329 [please put figure 6 around here]

330 [Please put figure 7 around here]

331

332 The last step of the visual analysis focused on an arbitrary transect, chosen as the diagonal
333 line connecting the North and the South vertices of the experimental site. The spatial evolution
334 of each predicted parameter along this transect is depicted in **Figure 8**. This approach allows
335 examining simultaneously the variability of the different parameters, namely of EF, SSM, F_r ,
336 and T^* . The results of this analysis are depicted in **Figure 8**. As one can notice, the variability
337 of the predicted parameters within the field follows largely explainable trends, depending on
338 both F_r and T^* . This observation provided further evidence of the technique’s ability to
339 satisfactorily predict both EF and SSM in the field when implemented with the UAV data.

340 [Please put figure 8 around here]

341

342 4.2 Point Comparisons

343 The results which concerned point-wise (i.e. pixel level) comparisons are summarised in
344 **Table 2**. As already noted, ground measurements of the radiation and turbulent fluxes were
345 acquired at a single location within the experimental field. On the other, SSM measurements
346 were conducted at a total of eight sites across the field, in which two different irrigation

347 strategies were applied since 1 July 2019. In particular, sites 1 to 4 were maintained under full
348 irrigation, whereas sites 5 to 8 under water deficit conditions.

349 As can be observed (in **Table 2**), the “simplified triangle” achieved very good predictions of
350 both EF and SSM, which are in close agreement to the field observations and in the same
351 range as the results of similar studies (e.g., Peng and Loew, 2014; Bai et al., 2019). The
352 predicted EF value, compared with the observed one, was slightly overestimated, with an
353 absolute difference of 0.053. However, it should be noted that this difference is also based on
354 a single ground measurement, since there was only one eddy covariance station installed in
355 the central part of the experimental site. In reference to the soil water content, Table 2 shows
356 that the predicted SSM is in very good agreement with the respective measurements, with
357 RMSE of 0.040 cm cm⁻³. Scatter (0.031 cm cm⁻³) contributes to RMSE relatively more than
358 Bias (-0.025 cm cm⁻³) but not overly so.

359 [please put Table 2 around here]

360

361

362 As shown in **Table 2**, the mean predicted SSM (denoted as “P”) for the locations of Stations 1
363 to 4 (plots with full irrigation) is 0.123 cm cm⁻³ while for locations of stations 5 to 8 (plots with
364 deficit irrigation) the mean predicted SSM is lower at 0.096 cm cm⁻³. On the other hand, the
365 measured SSM by the stations (denoted as “O”) does not reveal remarkable differences
366 between plots maintained under different irrigation strategies. The mean observed SSM for
367 plots 1 to 4 is 0.138 cm³ cm⁻³, while for the plots 5 to 8 it is only marginally lower and equal to
368 0.131 cm cm⁻³. While bias is generally low, the predicted SSM underestimates the
369 corresponding values in all the plots under deficit irrigation by -0.035 cm cm⁻³ on average.
370 However, for the fully irrigated plots, the underestimation is less than half in magnitude (equal
371 to -0.015 cm cm⁻³). All in all, these results suggest that the “simplified triangle” performed
372 satisfactorily in predicting both the EF and SSM under the examined conditions.

373

374 **5. Discussion**

375 Based on the results obtained (Section 4), the “simplified triangle” technique performed well to
376 in reproducing the high spatial resolution of EF and M_o /SSM maps for the study area. Both
377 predicted maps exhibited a largely explainable spatial variability across the experimental site,
378 with patterns in agreement to land cover type, topography and other site-specific
379 characteristics. In terms of statistical agreement, prediction accuracy was good for both EF
380 and SSM, and in agreement to the accuracies reported by other independent investigators
381 using different approaches and EO data types. For EF the difference between the predicted
382 and measured value is 0.053, giving a slight overestimation. After the M_o was converted to
383 SSM for the 8 stations, the results showed fairly low RMSE (0.040 cm cm⁻³) and low
384 underestimation (Bias = -0.025). These values are close to those reported by other studies
385 retrieving EF and SSM using TIR-based techniques (e.g., Peng and Loew, 2014; Nutini et al.,
386 2014; Lu et al., 2015; Xu et al., 2018; Bai et al., 2019). Thus, findings, although are based on
387 the single image analysis, are confirming the usefulness of the examined technique for EF and
388 SSM spatial determination at very fine resolution when implemented with UAV data.

389 There are a few factors which should be taken into consideration as well, when interpreting
390 the statistical agreement found herein. For example, the accuracy of the retrieved F_r and of T_s
391 is a possible cause of error as the technique requires only those two parameters as inputs for
392 its implementation. In our study, LST was measured by FLIR SC660 with an error lower than
393 0.03 °K, which is considered very small. Furthermore, since T_s is scaled in the “triangle”, the
394 effect of the predicted temperature accuracy might be small (Carlson, 2007). Possible reasons
395 for the lack of complete agreement could be related to the scale-mismatch between the EO-
396 data and the in-situ measurements, geo-location errors, and surface heterogeneity at the UAV
397 sensor spatial resolution, even though in this particular case predictions were obtained at very
398 high spatial resolution. Another possible factor concerning the SSM comparisons in particular
399 is that the ground measurements were acquired at 0 to 10 cm depth, while the UAV-derived
400 ones respond to soil water content at a much shallower layer (0 to 5 cm) over bare soil.
401 Effective soil depth for SSM measurement is an issue under investigation (Amani et al., 2016).
402 Some studies (Finn et al., 2011; Kasim et al., 2020) suggest an effective measurement at a
403 depth of 5 cm, while other studies (Zhang et al., 2015) suggest effective agreement at a depth
404 of 10 cm. Furthermore, uncertainties due to the instrumentation accuracy for EF and R_n
405 measurement should further be considered. Various studies have reported that errors in
406 instantaneous LE flux measurement can be in the order of 20% to 30%, which can be even
407 higher under certain circumstances (such as terrain features); similarly a measurement
408 uncertainty for R_n of 10% is not uncommon (Petropoulos et al., 2013).

409 Despite the promising results obtained in this first investigation performed herein, the
410 “simplified triangle” technique has some limitations which should also be acknowledged.
411 Those include its requirement to have within the image field of view a sufficient variability of F_r
412 and SSM range, in order to properly define the “wet” edge and the “dry” edge. Another issue is
413 the possible human error in the selection of warm and cold edges. However, this is an issue
414 common to other T_s/VI methods (Tomas et al., 2014; Mi et al., 2015). Furthermore, the
415 technique assumes a linear relationship between the TS/VI feature space and the predicted
416 EF and SSM, which might not necessarily be the case in nature.

417 Nonetheless, the “simplified triangle” capitalises on the inherent relationships existing in the
418 T_s/VI feature space for estimating M_o and EF. Yet, it seems to have some strong advantages
419 in comparison to other T_s/VI methods. The technique is simple to be applied and is dependent
420 on a few input parameters which can be easily computed from EO sensors. This makes the
421 technique implementation quick and computationally inexpensive when that is to be applied to
422 small scale studies. Its implementation, particularly with UAV images, presents several
423 advantages. When the technique is implemented with UAV data cloud cover is not an issue
424 (as UAVs fly at very low altitude) as it would be if satellite data had been used. In addition, the
425 technique when implemented with UAV data, the spatiotemporal variability of EF and M_o are
426 computed at a very fine spatial resolution (at 4 cm in our case). As information on very high
427 spatial and potentially temporal resolution of EF and SSM is essential to decision making in
428 most agricultural applications, including precision agriculture (Wang et al., 2018; Cui et al.,
429 2020), the potential added value of the “simplified” triangle technique to addressing this
430 requirement is clear. In overall, all the above characteristics place the “simplified triangle” in a
431 privileged position as a candidate for further investigation for a potential operationalisation with
432 either with satellite or airborne EO data.

433

434 **6. Conclusions**

435 In this study, a first assessment of the so-called “simplified triangle” technique was performed
436 to evaluate the ability of this method to predict EF and M_o/SSM when very high spatial
437 resolution EO imagery acquired from UAV are available. A robust evaluation was carried out
438 for an experimental site located in Sicily, Italy for which an extensive field campaign took place
439 in July 2019. To our knowledge, the study represents the first detailed assessment of this
440 innovative method with UAV data, particularly in a Mediterranean setting. The implementation
441 of the investigated herein technique with UAV images presents several advantages. Data
442 cloud cover is not an issue for UAV images and the spatiotemporal variability of EF and
443 M_o/SSM are computed at a very fine spatial resolution (at 4 cm in our case). Regardless, UAV
444 images present an additional challenge in correctly implementing the “simplified triangle”
445 technique. The method requires a sufficient variability of F_r and SSM range within the image
446 which can prove challenging in UAV imagery.

447 The obtained results suggest that the “simplified triangle” performed satisfactorily in predicting
448 both the M_o/SSM and EF. Validation showed an average error of 0.053 for EF and of 0.040
449 $\text{cm}^3 \text{cm}^{-3}$ for SSM. This implies that the investigated approach performs well under the semi-
450 arid conditions characterizing the experimental set up. Both predicted maps also exhibited
451 sensible spatial variability across the experimental site, with patterns in agreement to land
452 cover type, topography and other site-specific characteristics. The prediction accuracy of the
453 technique was also in close agreement, or even better, than accuracies reported by other
454 independent investigators using different T_s/VI approaches and EO data types.

455 However, the results reported herein are evidently based on a single image analysis. As the
456 technique is recent, further scrutiny and additional studies are required to establish its
457 applicability to different ecosystems. Such future investigations would require exploring the
458 prediction accuracy of the technique in different ecosystem environments and for longer time
459 periods using UAV imagery and spaceborne datasets from appropriate sensors (e.g. Landsat,
460 Setinel 1 to 3, Moderate Resolution Imaging Spectroradiometer (MODIS)), as well as including
461 a flux footprint analysis comparisons for the case of EF/ET predictions. In addition, a detailed
462 sensitivity analysis of the method would also allow quantifying the effect of T_s and F_r errors on
463 prediction accuracy. Other aspects of the technique that deserved investigation involve
464 automating the process of determining the wet and dry edge, which would also eliminate user
465 subjectivity in the technique implementation. It could potentially prove beneficial to combine
466 pixels to satellite sensor spatial resolution (e.g. from the Landsat resolution of 30 or 120 m) to
467 define the triangle boundaries. Then, once those boundaries have been established, they
468 could be imposed on the higher resolution UAV image. All the above are topics of key
469 importance that will be pursued in future studies.

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478

479 **Author contributions**

480 AM, GP, GC, and SM conceived and planned the experiments. AM and FC contributed to the
481 experimental design of the spectroradiometric acquisitions and radiometric calibration of the images.
482 GC, GP and SM coordinated the experiment and provided instrumentations. AM and GD
483 contributed to the GNSS experimental design and processing. GP contributed to the experimental
484 design and management of the soil moisture probes and processed the data. AM processed the flux
485 tower data. SM and FC designed and acquired the UAV images. GPP, AP, TNC, DH and CC
486 contributed to model implementation, results processing and analysis. GPP, AP prepared the
487 original draft of the manuscript. All authors reviewed and edited the final version of the
488 manuscript and contributed to the preparation of the revised manuscript.

489

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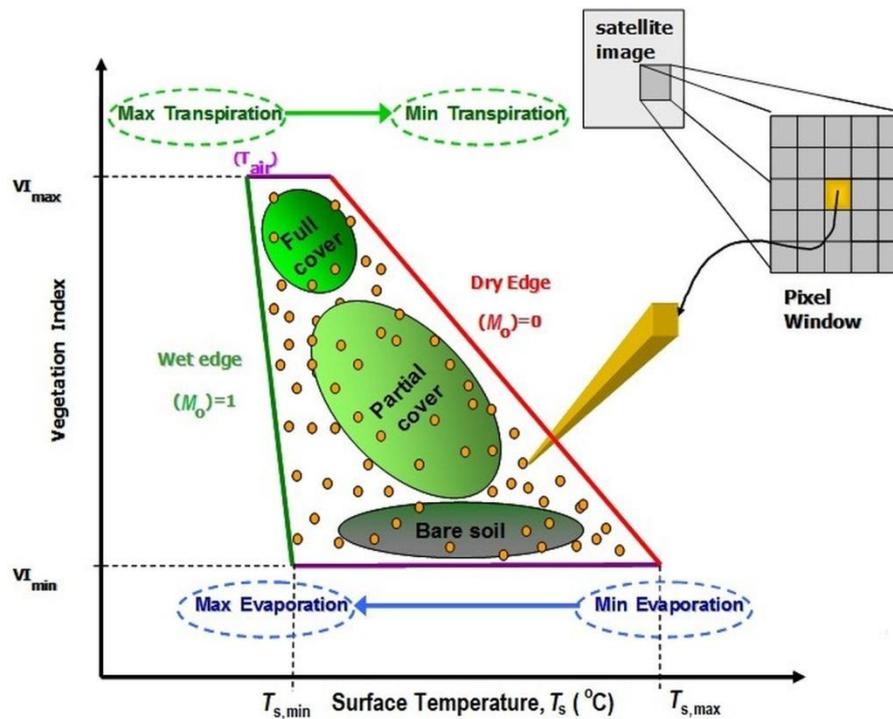


Figure 1: Conceptualisation of the main properties encapsulated in a T_s/VI scatterplot (adopted from Petropoulos et al. 2009).

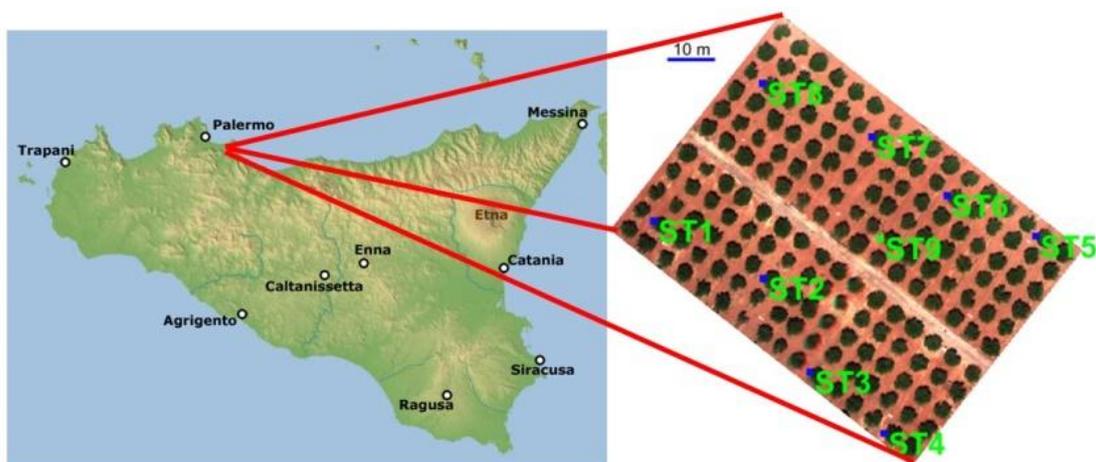


Figure 2: The experimental site, including the distribution of the ground measurement stations. ST1-8 refers to the locations of the probes that monitor soil water content, whereas ST9 is the eddy covariance system location. The image on the right is the actual UAV area covered by the UAV upon completion of orthorectification (see section 2.2.2 below).

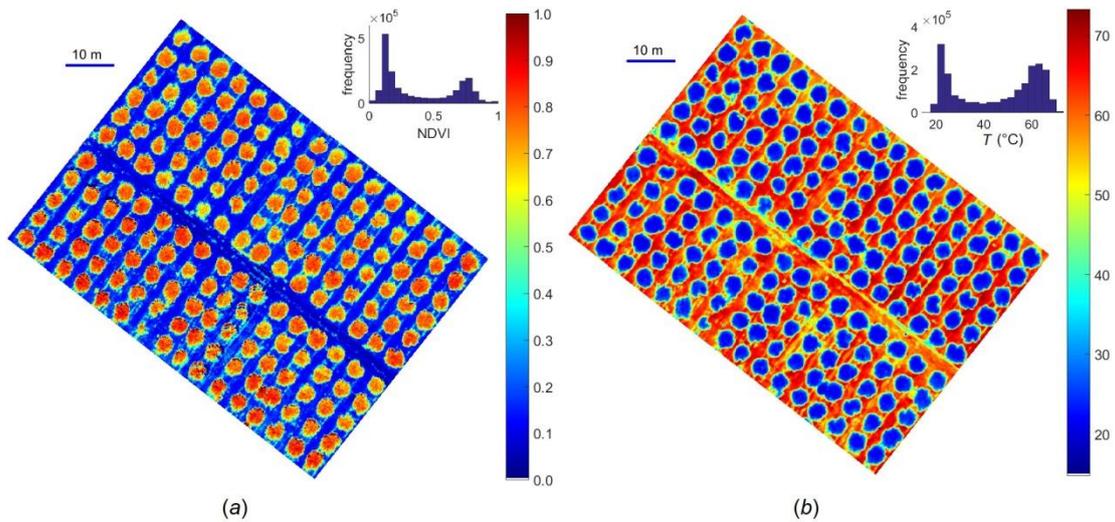


Figure 3: Pseudo colour maps of NDVI (a) and T_s ($^{\circ}\text{C}$) (b) derived upon completion of the pre-processing steps. The insets show the frequency histograms of NDVI and T_s respectively. Temperature units are in Celsius.

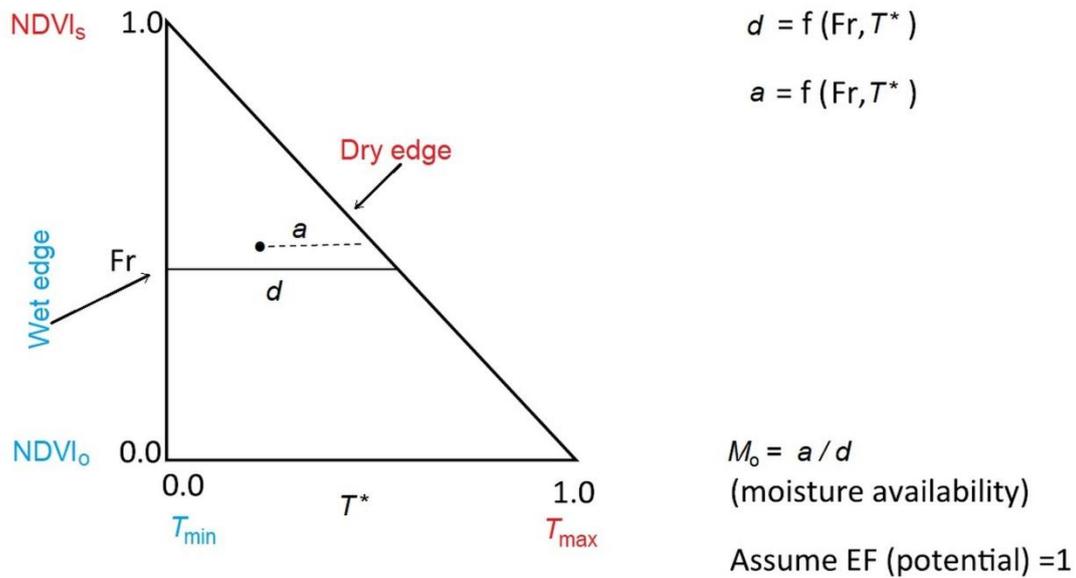


Figure 4: Graphical summary of the “simplified” triangle method principles and critical points selection required in its implementation (adopted from Carlson & Petropoulos, 2019)

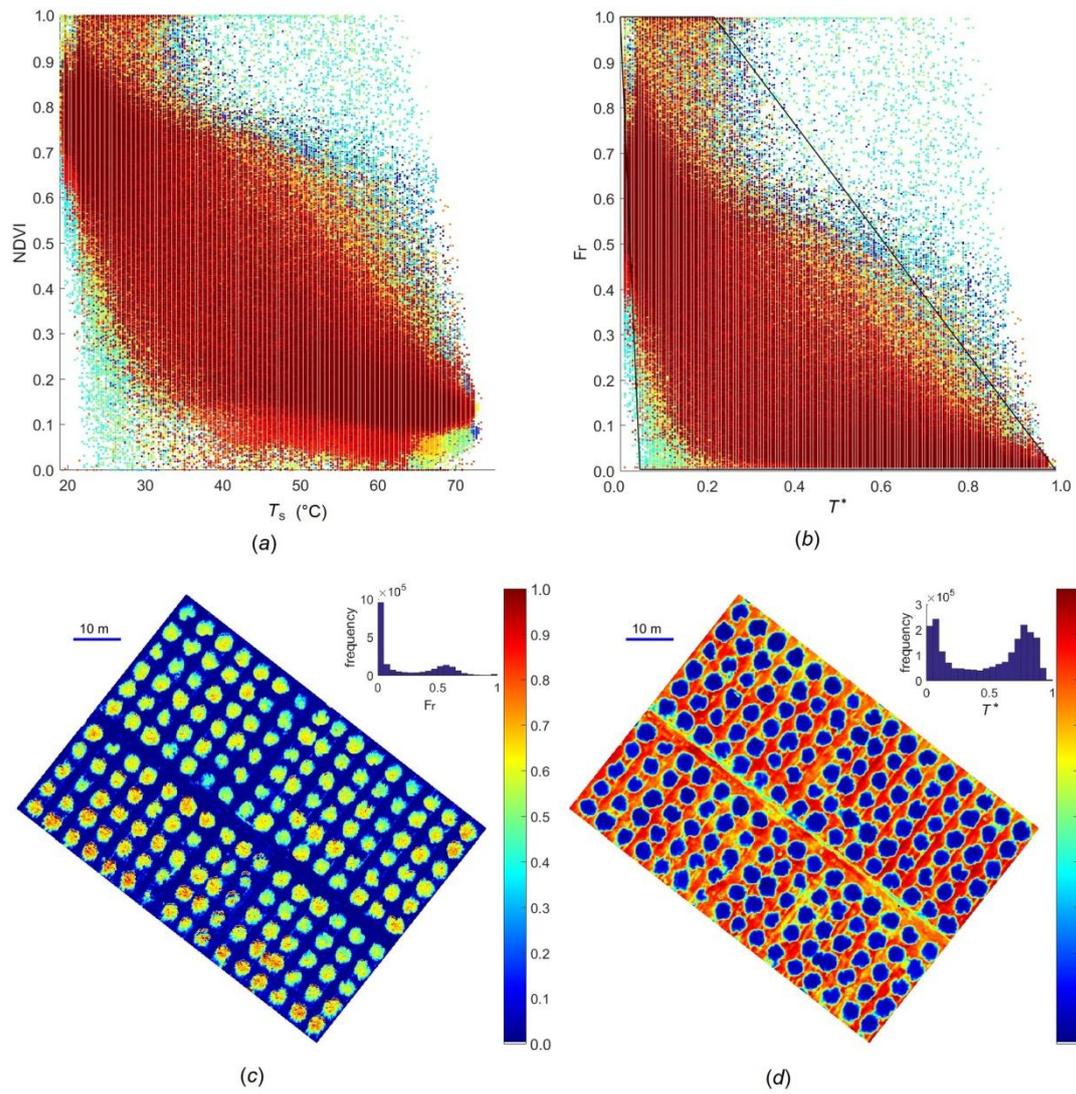


Figure 5: The scatterplots derived during the implementation (a,b), the F_r map (c) and the T^* map (d), derived from the datasets acquired with UAV. The “wet” and “dry” edge of the proposed triangle is shown by the continuous black line in scatterplot (b). The different colors in scatterplots (a,b) are for illustrative purposes only.

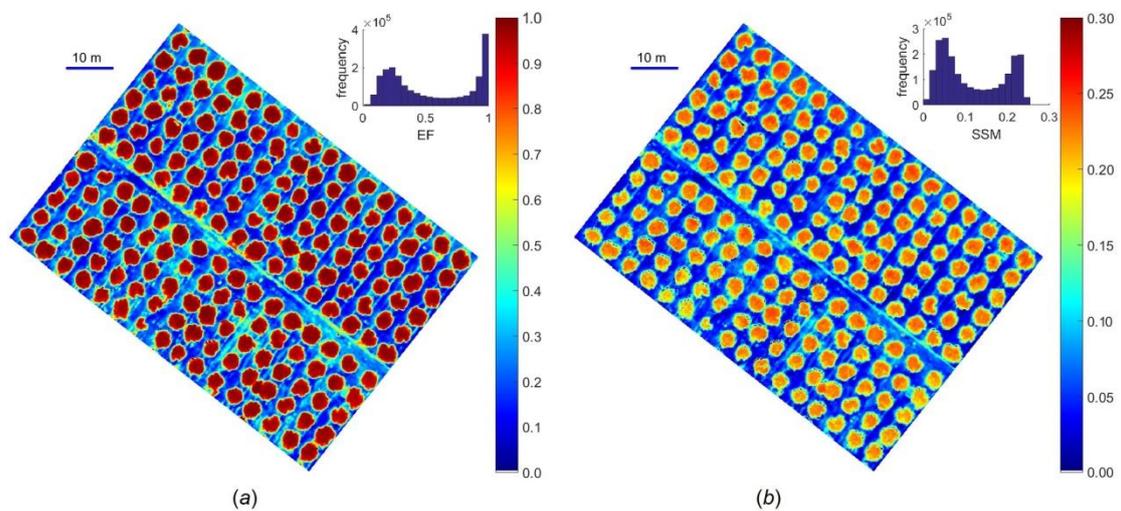


Figure 6: Maps of EF (a) and SSM (b) computed from the “simplified triangle” implementation using the data retrieved with UAV. The corresponding histograms are also shown.

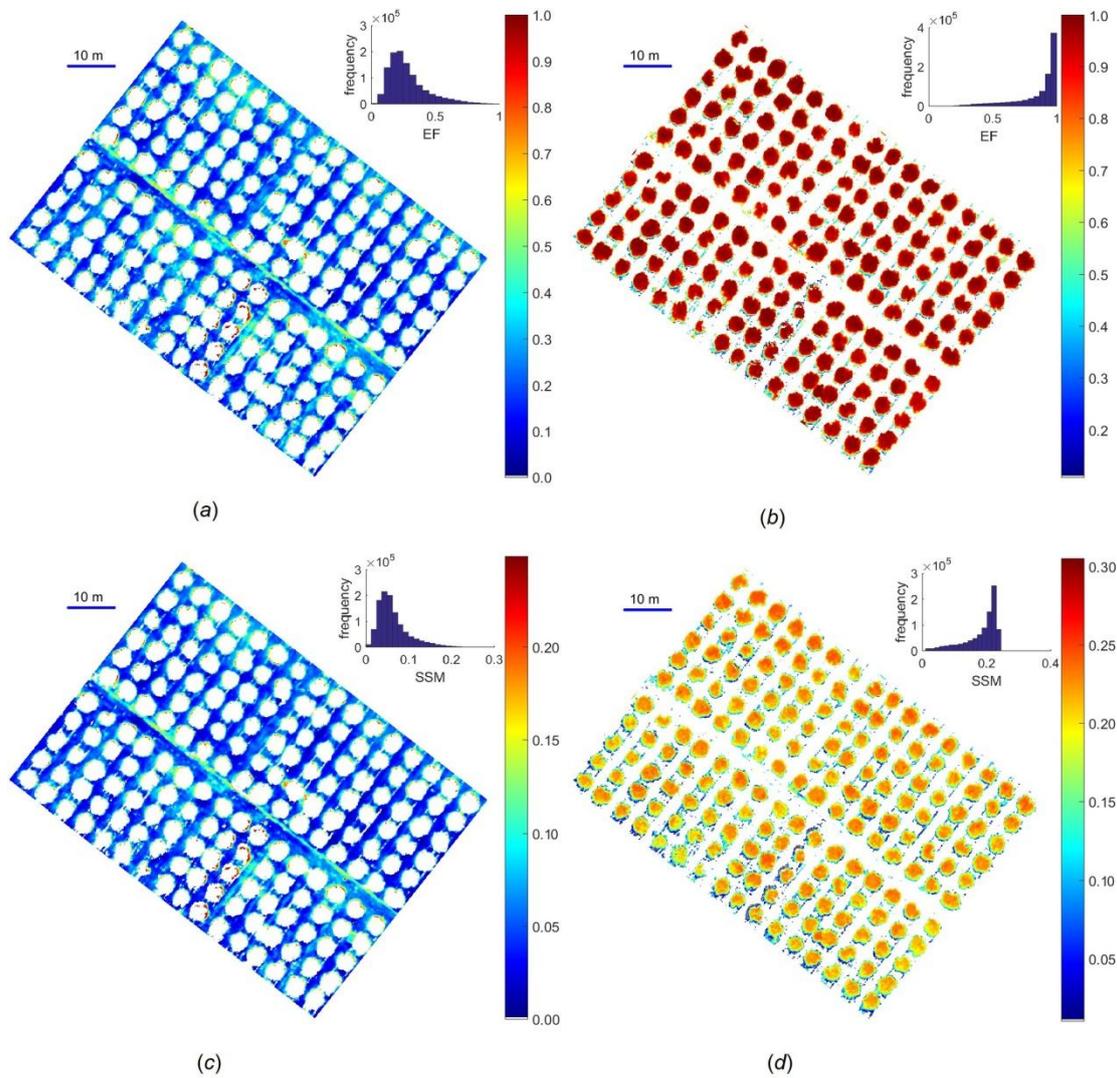


Figure 7: EF maps computed separately for the vegetated area (b) and for bare soil (a). Similarly, the derived SSM maps for the vegetated area (d) and for the bare soil (c), are also shown. Each map is accompanied by the corresponding frequency histogram.

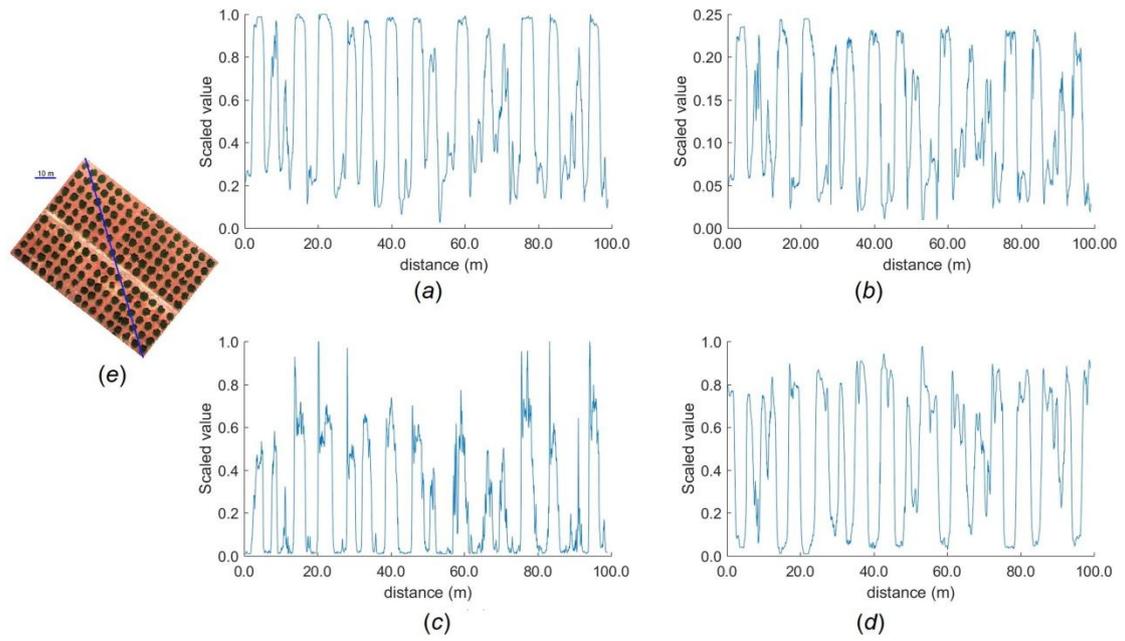


Figure 8: Arbitrarily selected transect within the field (e), and plots of the spatial variation of EF (a), SSM (b), F_r (c) and T^* (d) along the selected transect.

Table 1: Statistical measures used to assess the agreement between the predictions and ground observations. Subscripts $i = 1 \dots N$ refer to the individual observations, while O and P refer to the observed and predicted values.

Name	Description	Mathematical definition
Bias / MBE	Bias (accuracy) or Mean Bias Error	$\text{Bias} = \sum_{i=1}^N \frac{P_i - O_i}{N}$
Scatter / SD	Scatter (precision) or Standard Deviation	$S = \sqrt{\sum_{i=1}^N \frac{(P_i - O_i - (\overline{P_i - O_i}))^2}{N}}$
RMSE	Root Mean Square Error	$\text{RMSE} = \sqrt{\frac{\sum (P_i - O_i)^2}{N}}$

Table 2: Summary of the point by point comparisons between the ground observations (O) and the corresponding predicted with the “simplified triangle” (P). The differences (D) between predicted and observed values are also indicated. Bias, Scatter and RMSE are expressed in units of $\text{cm}^3 \text{cm}^{-3}$.

Fluxes (-)	O	P	D
LE R_n^{-1}	0.266	0.319	0.053
SSM ($\text{cm}^3 \text{cm}^{-3}$)	O	P	D
SM1	0.139	0.090	-0.049
SM2	0.107	0.132	0.025
SM3	0.162	0.171	0.009
SM4	0.145	0.099	-0.045
SM5	0.078	0.073	-0.006
SM6	0.121	0.084	-0.037
SM7	0.145	0.084	-0.061
SM8	0.180	0.144	-0.036
		Bias	-0.025
		Scatter	0.031
		RMSE	0.040